

# Miocene climate change: A molecular approach

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Alkenones are diagenetically resistant lipids known to be produced by only coccolithophorid Haptophyte algae (for example, *Emiliania huxleyi* today). These compounds are present ubiquitously in marine sediments of at least Miocene to Recent age, but are known from sediments as old as Cretaceous in age. The alkenones can be readily isolated from the organic fraction of sediment samples from DSDP/ODP sites. Field and experimental data have established that fractionation of carbon isotopes during photosynthesis (carbon fixation) for the Haptophyte algae is primarily controlled by the concentration of ambient  $\text{CO}_2(\text{aq})$  (a function of  $p\text{CO}_2$ , the partial pressure of  $\text{CO}_2$  in the atmosphere and/or surface- water mass) and cellular growth rate. These relationships have been calibrated, relating  $\epsilon_p$  (the difference between the carbon isotopic composition of the organic compound and the inorganic carbon source) to  $[\text{CO}_2(\text{aq})]$  and phosphate concentration in surface waters. This technique constitutes a paleo- $p\text{CO}_2$  barometer, provided that the  $\delta^{13}\text{C}$  of alkenones and dissolved inorganic carbon (DIC) can be measured, directly or indirectly, and growth rates estimated using various paleochemical proxies. Fortunately, a close

approximation of  $\delta^{13}\text{C}_{(\text{CO}_2\text{aq})}$  can be made by measuring the carbon isotopic composition of coeval planktic foraminifera and assuming equilibrium conditions among the carbon species in surface waters; required temperature estimates are derived from the oxygen isotopic compositions of selected planktonic foraminifers.

Using these techniques, we have been developing a detailed record of the  $\delta^{13}\text{C}$  of di-unsaturated alkenones for the early through middle Miocene from sequences recovered in DSDP/ODP sites in order to assess the role of postulated changes in  $p\text{CO}_2$  in climate change. The middle Miocene is a pivotal time of the Cenozoic, bridging times of fluctuating glacial conditions and with that of large, permanent continental ice sheets. Decreasing  $p\text{CO}_2$  and/or alterations in heat and vapor transport via rearrangement of deep-water circulation are commonly cited drivers of late Cenozoic global cooling. Our approach is to establish  $\epsilon_p$  records from low-productivity regions in the world's oceans, thereby mitigating the effect of variable growth rates on the  $\delta^{13}\text{C}$  of alkenones. As an example of the approach,

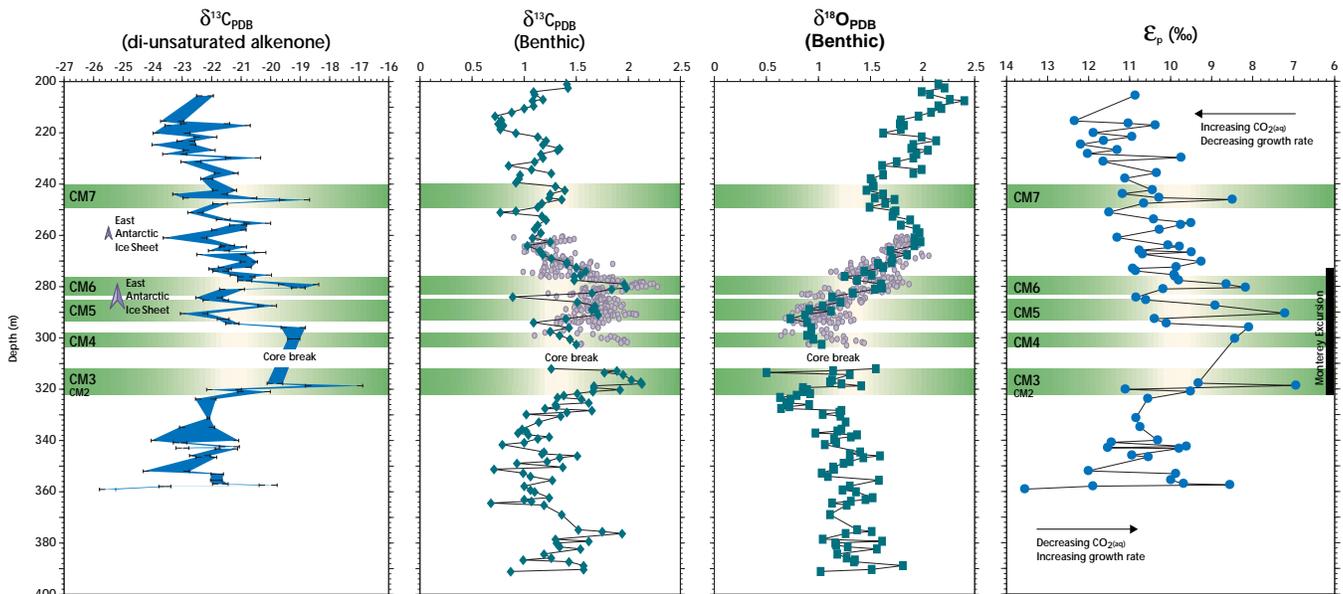


Figure 1A: Carbon isotopic composition of di-unsaturated alkenone from Site 588 in the southwest Pacific. Green banding indicates the range of  $\delta^{13}\text{C}$  maxima (CM2-CM7) within and outside the "Monterey Excursion."

Figure 1B: Carbon isotopic composition of benthic foraminifera from Site 588 [Kennett, 1985; Flower and Kennett, 1993].

Figure 1C: Oxygen isotopic composition of benthic foraminifera from Site 588 [Kennett, 1985; Flower and Kennett, 1993].

Figure 1D:  $\epsilon_p$  record calculated from the  $\delta^{13}\text{C}$  of di-unsaturated alkenones (figure 1A) and the  $\delta^{13}\text{C}$  of coeval planktic foraminifera.

we present preliminary data and interpretation from DSDP Site 588 (figure 1) in the southwest Pacific ocean. Site 588 records a continuous Miocene section with a low organic content and a low, steady sedimentation rate.

Results for the interval from  $\approx 22$  to 8 Ma exhibit both long- and short-term trends in alkenone  $\epsilon_p$ . A sharp and rapid decline followed by an equally rapid rise in  $\epsilon_p$  at  $\approx 21.5$  Ma indicates a dramatic fall and return in  $[\text{CO}_2(\text{aq})]$  and inferred  $p\text{CO}_2$ . A trend toward lower  $\epsilon_p$  follows this event, reaching a minimum throughout the 'Monterey excursion' of the middle Miocene ( $\approx 17.5$ -13.5 Ma). The 'Monterey excursion' is a broad, 1‰ positive shift in the  $\delta^{13}\text{C}$  of benthic foraminifera attributed to increased organic-carbon burial, that resulted in lower  $p\text{CO}_2$  and global cooling [Vincent and Berger, 1985].

Sharp positive shifts in  $\epsilon_p$  coincide with inorganic  $\delta^{13}\text{C}$  carbon maxima (CM2-CM7) within and outside the "Monterey excursion". These carbon maxima are associated with positive shifts in benthic  $\delta^{18}\text{O}$  and increased carbonate preservation [Woodruff and Savin, 1991], again suggesting a relationship between episodes of enhanced organic-carbon burial, decreasing  $p\text{CO}_2$ , and deep-water cooling. Our results support such interpretations. Furthermore, a trend toward higher  $p\text{CO}_2$ , inferred from the increasing  $\epsilon_p$  values following East Antarctic ice sheet growth at  $\approx 14.5$  Ma [Flower and Kennett, 1993], may highlight the importance of ice albedo in the regulation of global temperatures. Once global albedo increases a new climate system equilibrium is reached, allowing ice sheets to persist at a higher  $p\text{CO}_2$  than that required to initiate global glaciation.